



The composition of zircon in Variscan granites from Northern Portugal

Composición del circón de granitos variscos del Norte de Portugal

H.C.B. Martins, J. Abreu

Geology Centre. Department of Geosciences Environment and Spatial Planning. Faculty of Sciences, Porto University. Rua do Campo Alegre 4169-007 Porto, Portugal. E-mail: hbrites@fc.up.pt

ABSTRACT

A group of slightly peraluminous Variscan plutons in Northern Portugal were selected from the study of zircon composition. The selected plutons are: the Vila Pouca de Aguiar and the Lavadores-Madalena plutons with I-type affinities and the Vieira do Minho pluton, an I-S transitional type. Zircon occurs as euhedral to subhedral crystals and exhibit finely concentric oscillatory magmatic zoning mainly related to variations of Hf, Y, U and Th concentrations. Most zircon crystals show the dominant "xenotime" substitution. The zircon crystals have Zr/Hf ratio in the range of 21 to 52, with no significant differences between the different granites. These values are in the same range of other peraluminous granites and are in accordance with a crustal signature of zircon. Moreover, the range of Zr/Hf values in zircon crystals overlaps with that of crustal sources and consequently to the potential protoliths proposed in the genesis of the Vieira do Minho and the Vila Pouca de Aguiar plutons, namely meta-igneous crustal sources at different levels. Although zircon from the Lavadores-Madalena pluton has a compositional range similar to the other plutons, an origin by hibridisation has been proposed. However, similar zircon chemistry between this pluton and Vila Pouca de Aguiar and Vieira do Minho plutons could also suggest a similar crustal source.

Keywords: zircon composition; trace elements; Variscan granites

RESUMEN

Se han seleccionado tres plutones graníticos variscos en el norte de Portugal para el estudio de la composición del circón. Los plutones son: Vila Pouca de Aguiar y Lavadores-Madalena con afinidad de tipo-I y el plutón de Vieira do Minho de tipo transicional I-S. Los circones se presentan en cristales euhédricos a subhédricos y tienen zonados magnáticos, concéntricos oscilatorios finos ligados principalmente a variaciones de las concentraciones del Hf, Y, U y Th. La mayoría de los cristales de circón muestran la sustitución dominante "xenotima". Los zircones tienen relaciones Zr/Hf que varían en el rango 21–52, sin diferencias significativas entre los diferentes granitos. Estos valores son idénticos a otros granitos peralumínicos y son consistentes con circones de origen cortical. Además, las relaciones Zr/Hf coinciden con las de protolitos corticales y, en consecuencia con los propuestos en la génesis de los plutones de Vila Pouca de Aguiar y Vieira de Minho, en particular rocas metaigneas. Aunque los circones del plutón Lavadores-Madalena tienen una composición similar a los otros plutones, fue propuesto un origen por hibridación. Sin embargo la similitud química de los circones de este plutón con los otros puede también sugerir un origen similar.

Palabras clave: composición de circón; elementos traza; granitos Variscos

Recibido el 31 de marzo de 2014 / Aceptado el 11 de noviembre de 2014 / Publicado online el 10 de diciembre de 2014

Citation / Cómo citar este artículo: H.C.B. Martins & J. Abreu (2014). The composition of zircon in Variscan granites from Northern Portugal. *Estudios Geológicos* 70(2): e018. <http://dx.doi.org/10.3989/egeol.41729.318>.

Copyright: © 2014 CSIC. This is an open-access article distributed under the terms of the Creative Commons Attribution-Non Commercial (by-nc) Spain 3.0 License.

Introduction

Zircon is an invaluable tool in countless geological studies due to its wide distribution in a spectrum of rock types and particularly in granites. The importance of this accessory mineral lies in the combination of its tendency to incorporate trace elements, its chemical and physical durability and its remarkable resistance to high-temperature diffusive re-equilibration (Watson, 1996; Watson & Cherniak, 1997). Although the abundance of zircon is low, it may strongly affect the behaviour of many trace elements during the crystallization of magmas, and understanding its compositional variation is thus important for investigating the evolution of magmatic silicic systems (e.g. Hoskin *et al.* 2000; Belousova *et al.*, 2002 and 2006; Claiborne *et al.*, 2010).

Zirconium and Hf have a nearly identical geochemical behaviour, and therefore most of the crust maintains near-chondritic Zr/Hf ratios of ~35–40 (Hoskin & Schaltegger 2003). As zircon is the primary reservoir for both Zr and Hf and preferentially incorporates Zr, crystallization of zircon controls Zr/Hf, imprinting low Zr/Hf on coexisting melt. Thus, low Zr/Hf is a unique fingerprint of effective magmatic fractionation in the crust. Age and compositional zonation in zircons themselves provide a record of the thermal and compositional histories of magmatic systems. High Hf (low Zr/Hf) in zircon zones demonstrates growth from fractionated melt (Claiborne *et al.*, 2006, 2010).

In addition to its dominant role in controlling zirconium and hafnium distribution during magma evolution, zircon may have a significant influence on the behaviour of rare earth elements (REE), Y, Th, U, Nb and Ta (Heaman *et al.*, 1990; Bea, 1996; Belousova *et al.*, 2002, 2006; Hoskin & Schaltegger, 2003).

Large ionic radii and high charges make these elements incompatible in many rock-forming silicate minerals and they generally become concentrated in the residual melts, where the eventual crystallisation of zircon is able to accommodate these elements. The abundance and ratios of these elements are potentially useful to distinguish zircons from different sources (Pupin, 2000; Belousova *et al.*, 2002; Pérez-Soba *et al.*, 2007).

In this paper, we study the composition of zircon from a group of late- to post- Variscan plutons in

NW of Portugal which present variable typology (Martins *et al.*, 2007, 2009 and 2013; Silva, 2010).

Geological and petrological features

The granitic plutons selected for this study are located the Central Iberian Zone and one of them on the western border of the same zone, NW Portugal, being the Vila Pouca de Aguiar pluton, the Vieira do Minho pluton and the Lavadores-Madalena pluton, respectively (Fig. 1). Their emplacement were controlled by tectonic regional structures during the late stages of the Variscan orogeny, mainly from 310 to 299 Ma (Martins *et al.*, 2007, 2009, 2011, 2013). The Vila pouca de Aguiar pluton consists of two main biotite-bearing granitic rocks of monzonitic composition, the coarse-grained Vila Pouca de Aguiar granite (VPAG) as the dominant facies and the more leucocratic medium-grained Pedras Salgadas granite (PSG). These two granites are slightly peraluminous ($0.99 < A/CNK < 1.11$) and have an I-type affinity. The U-Pb zircon dating yields a consistent age of 299 ± 3 Ma (Martins *et al.*, 2009). The Vieira do Minho pluton is composed by the porphyritic coarse-grained Vieira do Minho granite (VMG) and the medium-grained Moreira de Rei granite (MRG), both are biotite bearing granitic rocks of monzonitic-granodioritic composition (Almeida *et al.* 2002; Martins *et al.* 2013). They are peraluminous with A/CNK ranging from 1.04 to 1.24, and with a I-S transitional type affinity. The U-Pb isotopic analyses carried out on zircon and monazite from VMG indicate a crystallisation age of 310 ± 2 Ma (Martins *et al.*, 2013), and 308 ± 4 Ma for MRG (Dias *et al.*, 2002).

The Lavadores-Madalena pluton outcrops in a narrow band along the coast line, south of Porto river. The Lavadores granite (LG) is pinkish porphyritic, coarse-grained biotite granodiorite-monzogranite, with accessory amphibole. This granite presents abundant mafic microgranular enclaves with a wide range of compositions (Silva and Neiva, 1998; Silva, 2010). It has a gradual contact with the Madalena granite (MG), which show silimar characteristics. They are both weakly peraluminous ($0.98 < A/CNK < 1.04$) and have affinity with I-type granites. A U-Pb zircon age of 298 ± 11 Ma was reported for the Lavadores granite (Martins *et al.*, 2011).

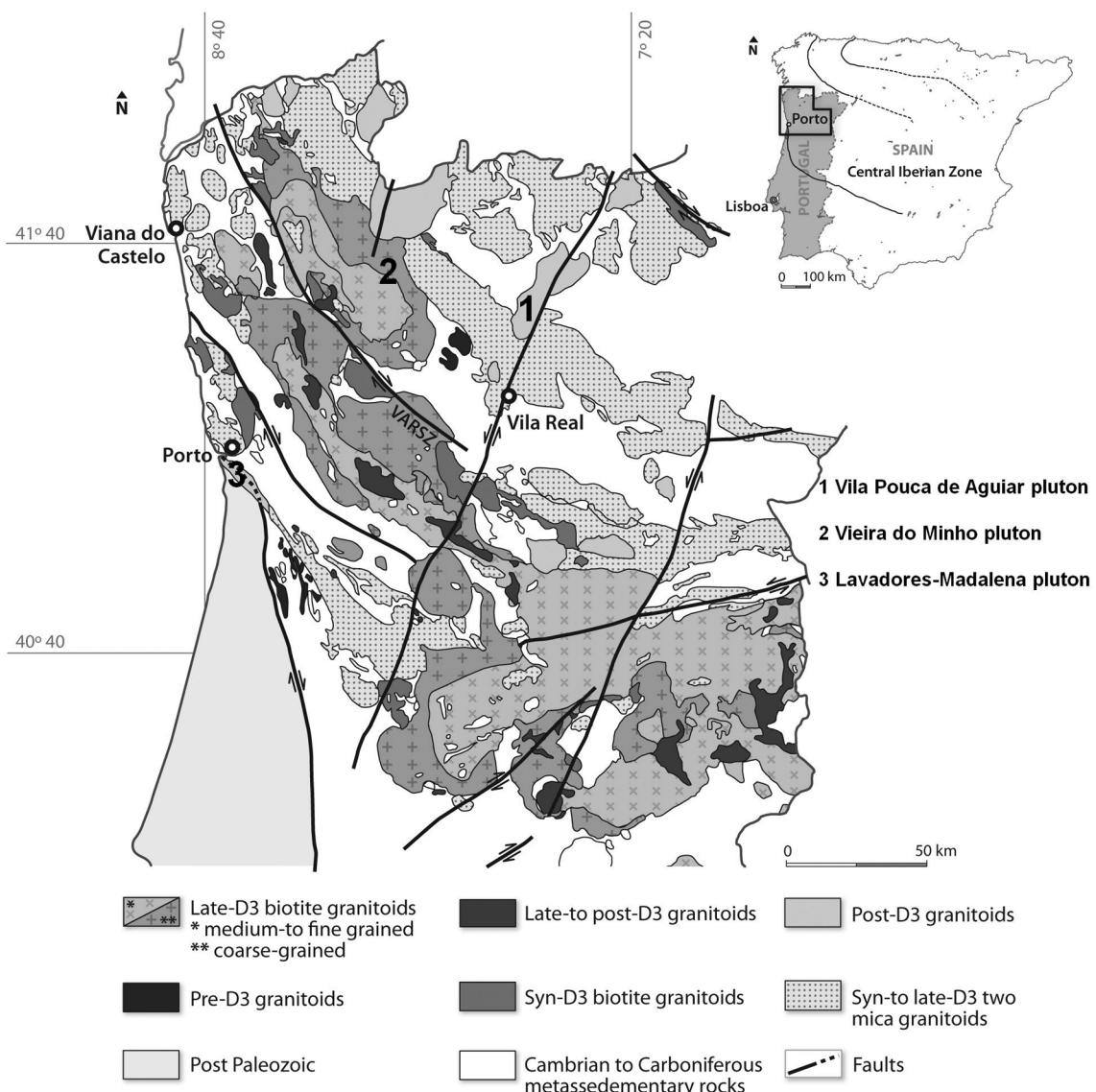


Fig. 1.—Geological distribution of Variscan syn- to post-orogenic granitoids in the Central Iberian Zone, northern Portugal (Ferreira *et al.* 1987, modified). 1, 2 and 3 locations of the studied plutons.

The mineralogical data obtained on these granites are summarized in Table 1.

Analytical methods

Zircon was separated from <250 µm sieved fractions using standard heavy liquids and magnetic techniques, handpicked and mounted in epoxy resin. Imaging and chemical analyses were performed using a JEOL 8500 F electron microprobe at LNEG (Porto). Prior to quantitative analysis, the grains were

imaged using a high-energy back-scattered electrons were recorded using a paired semiconductor detector, operating in compositional mode to obtain atomic number contrast (BSE). Most zircon analyses were carried out using a focused beam of 1 µm, a 20 nA beam current and with an acceleration voltage of 20 kv. A ZAF correction program was used.

The detection limits were 244 ppm for Si, 240 ppm for Fe, 224 ppm for Al, 243 ppm for P, 421 ppm for Hf, 1042 ppm for La, 111 ppm for Ca, 663 ppm for

Table 1.—Summary of magmatic mineralogy of the studied plutons

Pluton		Texture	Major minerals	Accessory minerals
Vieira do Minho	monzogranite	porphyritic, coarse-grained	Qtz, Pl, Kfs, Bt	Ms, Zr, Mnz, Ap, Ilm, Crd, And
	granodiorite monzogranite	porphyritic, medium- to fine-grained	Qtz, Pl, Kfs, Bt	Ms, Zr, Ap, Ilm, Aln, Thr
Vila Pouca de Aguiar	monzogranite	porphyritic, coarse-grained	Qtz, Pl, Kfs, Bt	Zr, Ap, Aln, Ilm
	monzogranite	porphyritic, medium-grained	Qtz, Pl, Kfs, Bt	Ms, Zr, Ap, Mnz, Xnt, Aln, Ilm
Lavadores-Madalena	granodiorite monzogranite	porphyritic, coarse-grained	Qtz, Pl, Kfs, Bt	Hbl, Zr, Ap, Mnz, Aln, Mt
	granodiorite monzogranite	porphyritic, medium-grained	Qtz, Pl, Kfs, Bt	Zr, Ap, Mnz, Ilm, Mt

Pr, 836 ppm for Ce, 479 ppm for Zr, 667 ppm for Nd, 222 ppm for Ti, 786 ppm for Gd, 514 ppm for Th, 156 ppm for U and 421 ppm for Y.

zoning exhibits a compositional trend of decreasing ThO_2 and Y_2O_3 and increasing UO_2 from core to rim.

Zircon Chemistry

Nearly 150 chemical analyses of zircon were carried out in the studied plutons. Representative results are shown in Table 2. The zircon crystals are euhedral to subhedral and exhibit finely concentric oscillatory zoning, typical of primary zircon growth (e.g., Smith *et al.*, 1991; Hanchar & Miller, 1993). The zoning is characterized by anti-correlated BSE intensities. In general these crystals show a weakly zoned core surrounded by an outer oscillatory rim (Martins *et al.*, 2014). Internal zoning is mainly related to variations of Hf, Y, U and Th concentrations, so these elements can be main indicators of the degree of magmatic evolution. The compositional range of the elements analysed according to the studied plutons is presented in Figure 2. The compositional range is independent of their host granite, as described in other works (Pérez-Soba *et al.*, 2007 and references therein). It can be seen that all granites present similar medians and low values in the oxides considered, although with different ranges especially in UO_2 and Y_2O_3 (Abreu, 2012).

The Vila Pouca de Aguiar pluton

The compositional ranges in the Vila Pouca de Aguiar granite (VPAG) zircon are (in wt%): 1.29–2.07 HfO_2 , 0–0.07 UO_2 , 0–0.09 ThO_2 and 0–0.1 Y_2O_3 , whereas the Pedras Salgadas granite (PSG) zircon displays a slightly lower range in HfO_2 (1.24–1.9 wt%) and higher range in UO_2 (0–0.41 wt%), ThO_2 (0–1.28 wt%) and Y_2O_3 (0–4.22 wt%) (Fig. 2). In both granites the zircon

The Vieira do Minho Pluton

The Vieira do Minho granite (VMG) presents HfO_2 values ranging from 1.21 to 2.17 wt%, Y_2O_3 from 0 to 3.2 wt%, ThO_2 from 0 to 0.17 wt%. As regards the Moreira de Rei granite (MRG) the HfO_2 values are between 0.18 to 2.55 wt%, Y_2O_3 values from 0 to 0.57 wt% and the contents of ThO_2 range from 0 and 0.13 wt%. The UO_2 contents in both granites are low and very similar, showing values from 0 to 0.07 wt%. In both granites zoning shows a compositional trend of increasing HfO_2 , Y_2O_3 and UO_2 from core to rim whereas the contents of ThO_2 are in general distributed uniformly in the rim and core, but in some grains it also increase from core to rim.

The Lavadores-Madalena Pluton

The Madalena granite (MG) has the largest compositional range of HfO_2 (1.29 to 2.92 wt%) and UO_2 (0–1.8 wt%) compared to the Lavadores granite (LG) (HfO_2 : 2.04 to 1.04 wt% and UO_2 : 0–1.2 wt%). In both granites the Y_2O_3 and ThO_2 content are low and similar, being the highest value 0.5 wt% and 0.3 wt%, respectively. In general, it is observed in zircon of these granites a positive correlation of, U, Th and Y, which increases from core to rim. The HfO_2 contents from zircons of the Madalena granite increase from core to the rims. As we found in the Vila Pouca de Aguiar pluton, the HfO_2 contents of zircon from the Lavadores-granite are also distributed in an almost uniform way both in the core and in the rims in most of the analysed zircons.

Table 2.—Representative electron microprobe analyses of zircon in Variscan granites from Northern Portugal

Plutons	Vila Pouca de Aguiar												Vieira do Minho												Lavadores Madalena											
	VPAG						PSG						VMG						MRG						LG						MG					
	Z1	Z1	Z2	Z2	Z1	Z1	Z3	Z3	Z1	Z1	Z5	Z5	Z1	Z1	Z5	Z5	Z3	Z3	Z7	Z7	Z2	Z2	Z3	Z3	Z1	Z1	Z5	Z5	Z3	Z3	Z7	Z7	Z2	Z2	Z3	Z3
Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim			
SiO ₂	34.75	34.57	34.60	34.28	33.96	33.65	34.15	28.39	33.21	33.47	34.26	33.66	34.54	34.53	34.36	33.66	34.21	34.48	34.89	34.8	33.41	33.15	33.59	32.99												
ZrO ₂	61.56	61.42	63.36	63.5	64.36	63.33	64.7	64.34	63.17	62.79	63.74	63.13	64.84	64.35	63.79	63.65	63.45	63.37	63.96	62.85	65.07	64.33	64.68	64.57												
Al ₂ O ₃	0.03	0.06	—	—	0.1	—	0.02	—	0.04	—	0.08	—	0.01	—	—	—	0.01	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—			
P ₂ O ₅	0.18	0.15	0.06	0.24	0.16	0.62	0.09	0.11	0.22	0.21	0.17	0.36	0.14	0.18	0.11	0.41	0.11	0.01	0.06	0.07	0.02	0.07	0.08	—	—	—	—	—	—	—	—	—	—	—	—	
UO ₂	0.20	0.15	0.06	0.19	0.9	0.23	0.02	0.04	0.08	0.1	0.07	0.2	0.06	0.26	0.04	0.46	0.23	0.14	0.07	0.12	0.09	0.08	0.11	0.08	—	—	—	—	—	—	—	—	—	—	—	
Y ₂ O ₃	0.29	0.08	—	0.30	0.64	0.34	0.01	0.03	0.24	0.14	0.00	0.07	—	—	—	0.36	0.35	—	—	—	0.09	0.08	0.05	0.01	—	—	—	—	—	—	—	—	—	—	—	
HfO ₂	1.45	1.68	1.62	1.29	1.77	1.67	1.24	1.61	1.55	1.82	1.44	1.68	1.69	1.81	1.53	1.87	1.52	1.92	1.7	1.52	1.83	1.36	1.71	1.53	—	—	—	—	—	—	—	—	—	—	—	
ThO ₂	0.07	0.01	—	0.11	0.40	0.19	0.00	0.06	0.05	0.03	—	0.05	0.02	0.07	0.08	0.07	0.08	0.07	0.25	0.08	0.05	0.12	—	—	0.08	0.05	—	—	—	—	—	—	—	—	—	
La ₂ O ₃	—	0.05	0.03	—	—	0.54	0.12	0.09	—	0.03	0.03	—	—	0.07	—	0.16	—	0.23	—	—	—	—	—	0.00	—	—	—	—	—	—	—	—	—	—		
Pr ₂ O ₃	—	0.3	—	—	—	—	—	0.09	—	—	0	—	—	0.02	—	—	0.04	—	0.02	—	0.02	—	0.15	—	0.05	—	—	—	—	—	—	—	—	—		
Ce ₂ O ₃	—	—	—	0.06	—	0.04	—	0.07	0.19	—	0.09	—	0.08	—	—	—	—	—	—	0.10	—	—	—	0.03	—	0.1	—	—	—	—	—	—	—	—		
Nd ₂ O ₃	0.02	—	0.08	—	—	0.02	—	—	0.07	—	0.04	—	0.08	—	—	—	—	—	—	—	—	—	—	0.03	—	—	—	—	—	—	—	—	—	—		
Gd ₂ O ₃	—	—	—	—	—	—	—	—	0.03	0.07	—	—	—	—	—	—	—	—	—	—	—	—	—	0.02	—	—	—	—	—	—	—	—	—	—		
TiO ₂	—	—	0.04	—	—	—	0.01	—	—	0.01	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	0.02	—	—	—	—	—	—			
CaO	—	0.05	—	—	0.03	0.14	—	0.05	0.03	—	0.03	0.01	—	0.02	—	—	0.26	0.01	—	—	—	0.05	—	0	—	—	—	—	—	—	—	—	—	—		
FeO	0.04	0.07	—	0.06	0.01	0.03	0.04	0.02	—	0.05	0.05	0.12	—	—	—	—	0.02	0.02	—	—	0.02	0.03	0.03	0.02	—	—	—	—	—	—	—	—	—	—	—	—
Total	98.60	98.34	99.85	99.97	101.5	100.4	100.4	98.73	98.95	99.84	99.48	101.4	101.3	100.0	100.5	100.2	100.2	100.9	99.73	100.5	99.39	100.3	99.47	—	—	—	—	—	—	—	—	—	—	—	—	

horizontal line indicates contents below detection limit.

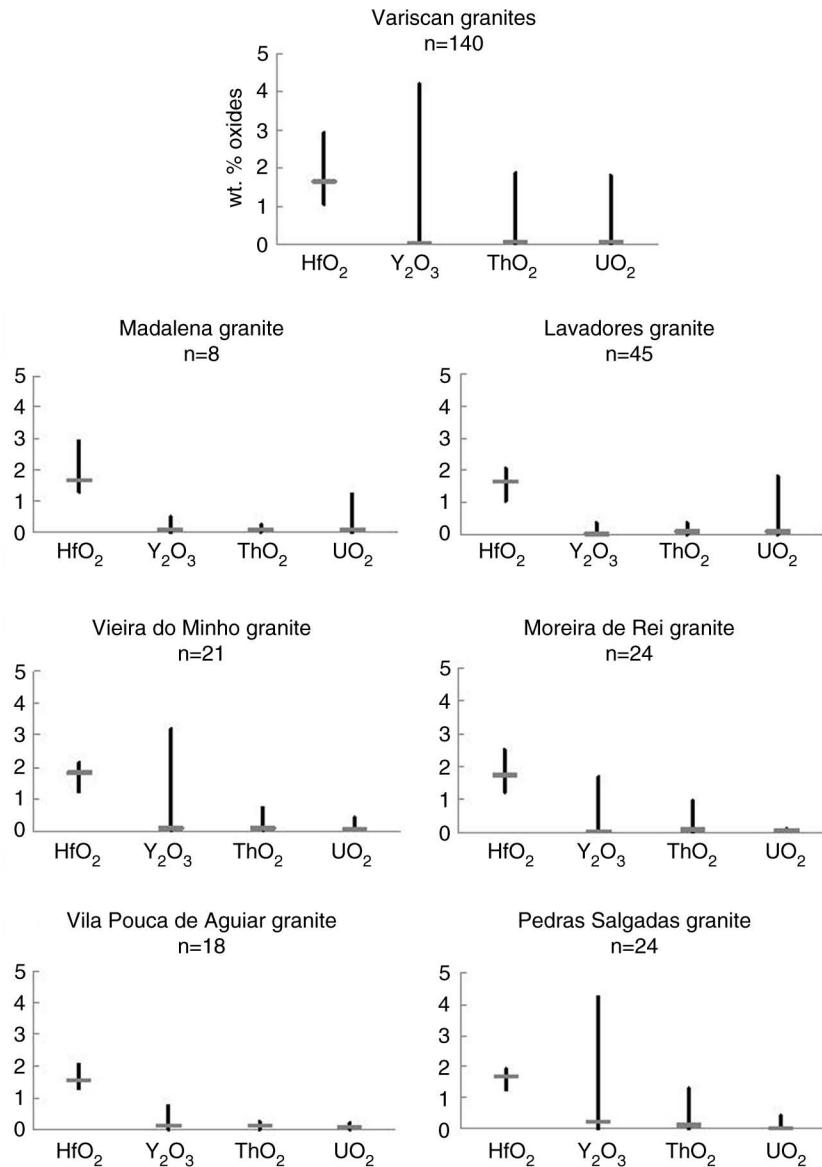


Fig. 2.—Compositional comparison between zircon of the studied granites. Bar shows the complete range of zircon composition (vertical line), including the median (short horizontal line).

Discussion and conclusions

The zircon compositional zoning observed in the studied granites is similar and common to other peraluminous granites, namely monzogranites and granodiorites from the Spanish Central System (Pérez-Soba *et al.*, 2007).

All zircon studied are large euhedral crystals and exhibit finely concentric oscillatory zoning in BSE images. These crystals commonly show an unzoned inner part and outer oscillatory bands.

Its composition is close to that of pure zircon, with increasing Hf, Y, U and Th contents from core to rim in granites from the Vieira do Minho Pluton and in the Lavadores granite, whereas in the Madalena granite and in the Vila Pouca de Aguiar pluton, zircons show a different trend with a decrease of Th, U, Y from core to rim and Hf shows an almost uniform distribution both in the core and in the rims in most of the analysed zircons. The decrease of Th-U-Y contents could suggest cocrystallization with other accessory phases rich in Y (xenotime), or rich in U and Th

(e.g. thorite, monazite and allanite). These accessory phases, except thorite, were found in these granites and may contribute to the decrease of the elements referred above.

Zr can be replaced by others tetravalent cations (Hf, Th, U) by single isovalent replacement or by coupled substitutions where divalent (Ca, Fe, Mg, Mn) and trivalent (REE, Y, Fe) cations may be incorporated (Hosking & Schaltegger, 2003). In the studied zircons, many elements behave coherently because they participate in a coupled substitution in the zircon crystal structure. Phosphorus concentrations show a positive correlation with Y content, reflecting the dominant “xenotime” substitution: $(\text{REE}, \text{Y})^{3+} + \text{P}^{5+} = \text{Zr}^{4+} + \text{Si}^{4+}$ (eg. Speer, 1980).

Uranium and Th substitute directly for Zr, however they both show a positive correlation with Y as well as with P. This widely reflects the trace element composition of the parental rocks, where the incompatible element concentrations tend to be higher since these rocks are fractionated granites.

The incorporation of Hf in the structure of zircon seems to be strongly dependent on temperature rather than the melt chemistry (Clairborne *et al.*, 2006; Watson *et al.*, 2006). On the other hand, the Th/U ratio has shown to be sensitive to temperature variations (e.g. Gagnevin *et al.*, 2010, and references therein). Nevertheless, in all the studied plutons, UO_2 and ThO_2 in zircon show no correlation with HfO_2 and Th/U ratio (Fig. 3). However, in both granites

of the Vieira do Minho Pluton and in the Madalena granite, the increasing Hf from core to rim in most of zircon grains suggests magma differentiation during crystal fractionation (Benisek & Finger, 1993; Hoskin & Schaltegger, 2003). Accordingly, in the Vila Pouca de Aguiar pluton and in the Lavadores granite, the homogeneous Hf concentrations from core to rim may suggest that zircon crystallization occurred in a short period before large fractionation of the magma which reflects the constant Zr/Hf ratio of the melt during zircon formation.

Because hafnium is nearly identical in size and charge to zirconium, the two behave nearly identically. Zircon, therefore, is essentially a zircon–hafnon solid solution, with most natural zircons containing between 1 and 2 wt% HfO_2 (Hoskin & Schaltegger, 2003; Belousova *et al.*, 2002; Bea *et al.*, 2006; Claiborne *et al.*, 2006 and references therein) and is the major reservoir for the crust’s Hf. As zircon is the primary reservoir for the crust’s Hf, growth of zircon controls the Hf composition of any melt. As zircon grows, both the Zr and Hf concentrations in the magma will decrease. However, this growth preferentially incorporates Zr over Hf and therefore leads to a decrease in the Zr/Hf ratio of the remaining melt. This process leads to increasing Hf concentration in growing zircon crystals, so the the Zr/Hf ratio has been used as an index of magmatic differentiation (Černy *et al.*, 1985). Nevertheless Pérez-Soba *et al.* (2007), suggest that Zr/Hf ratio can be used with caution as an index of magmatic fractionation, because Zr in zircon is also replaced by other elements and Hf may increase or decrease in a zoned single crystal of the same rock, which is observed in most of the zircons analysed.

The Zr/Hf ratio of the studied granites, ranges from 21 to 52, without significant differences between the tree plutons (mean Zr/Hf: VPAG=41, PSG=40; VMG=36, MRG=37; LG=39; MG=37) (Fig. 4). These values are in the same range observed in other peraluminous granites and are in accordance with a crustal signature of zircon, as suggested by several authors (e.g. Pupin, 2000; Pérez-Soba *et al.*, 2007). The range of Zr/Hf values in zircon crystals of the studied granite plutons overlaps with that of crustal sources and consequently to the potential protoliths proposed in the genesis of the Vieira do Minho and the Vila Pouca de Aguiar plutons (Martins *et al.*, 2009).

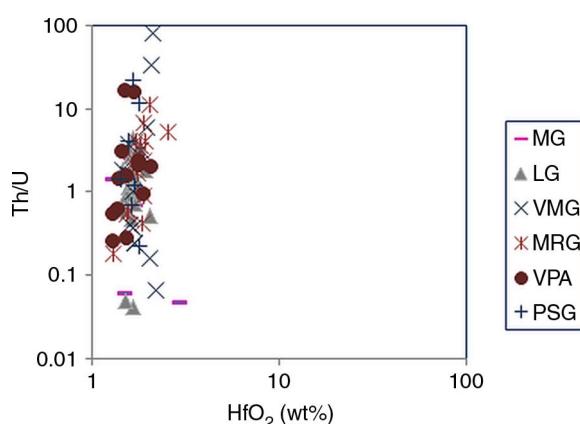


Fig. 3.—Co-variation of Th/U ratio with Hf of zircons from the studied granites. LG and MG (granites from the Lavadores-Madalena pluton); VMG and MRG (granites from the Vieira do Minho pluton); VPAG and PSG (granites from the Vila Pouca de Aguiar pluton).

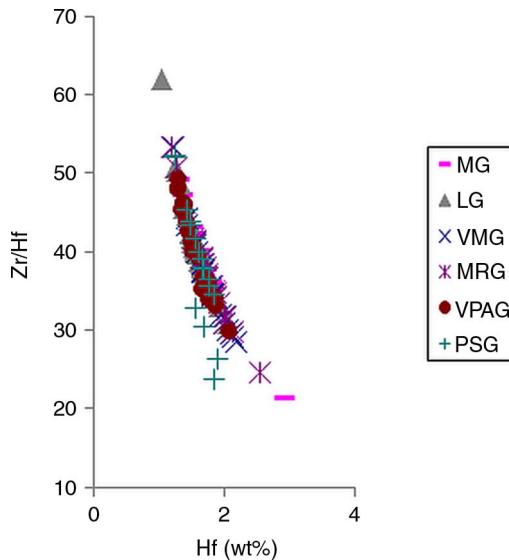


Fig. 4.—The composition of zircon in terms of Hf (wt%) versus Zr/Hf from the studied granites. LG and MG (granites from the Lavadores-Madalena pluton); VMG and MRG (granites from the Vieira do Minho pluton); VPAG and PSG (granites from the Vila Pouca de Aguiar pluton).

and 2013). Although zircon from the Lavadores-Madalena pluton has a compositional range similar to the other plutons, Silva (1995) has proposed an origin by hybridisation, between basic and felsic magmas, based on isotopic data. However, the similarity in zircon chemistry and also isotopic data between this pluton and Vila Pouca de Aguiar and Vieira do Minho plutons could also suggest an origin by metagranitic crustal sources.

ACKNOWLEDGEMENTS

This research was financially supported for the Geology Center (Porto University), an R and D unit (Unit 39) from Portuguese Foundation of Science and Technology. The authors would like to thank CEMUP (Porto University) for BSE images and LNEG for the microprobe zircon analyses. Helpful and constructive reviews by two anonymous reviewers are gratefully acknowledged.

References

- Abreu, J. (2012). Estudo geoquímico de zircão de granitos do NW Português. Master Thesis, Universidade do Porto, 176 pp.
- Almeida, A.; Martins, H.C.B. & Noronha, F. (2002). Hercynian acid magmatism and related mineralisations in Northern Portugal. *Gondwana Research*, 5: 423–434. [http://dx.doi.org/10.1016/S1342-937X\(05\)70733-4](http://dx.doi.org/10.1016/S1342-937X(05)70733-4).
- Bea, F. (1996). Residence of REE, Y, Th and U in Granites and Crustal Protoliths; Implications for the Chemistry of Crustal Melts. *Journal of Petrology*, 37: 521–552. <http://dx.doi.org/10.1093/petrology/37.3.521>.
- Bea, F.; Montero, P. & Ortega, M. (2006). A LA-ICPMS evaluation of Zr reservoirs in common crustal rocks: implications for Zr and Hf geochemistry and zircon-forming processes. *Canadian Mineralogist*, 44: 693–714. <http://dx.doi.org/10.2113/gscanmin.44.3.693>.
- Belousova, E.A.; Griffin, W.L.; O'Reilly, S.Y. & Fisher, N.I. (2002). Igneous zircon: trace element composition as an indicator of source rock type. *Contributions to Mineralogy and Petrology*, 143: 602–622. <http://dx.doi.org/10.1007/s00410-002-0364-7>.
- Belousova, E.A.; Griffin, W.L. & O'Reilly, S.Y. (2006). Zircon crystal morphology, trace element signatures and Hf isotope composition as a tool for petrogenetic modelling: examples from Eastern Australian granitoids. *Journal of Petrology*, 47: 329–353. <http://dx.doi.org/10.1093/petrology/egi077>.
- Benisek, A. & Finger, F. (1993). Factors controlling the development of prism faces in granite zircons: a microprobe study. *Contributions to Mineralogy and Petrology*, 114: 441–451. <http://dx.doi.org/10.1007/BF00321749>.
- Černý, P.; Meintzer, R.E. & Anderson, A.J. (1985). Extreme fractionation in rare-element granitic pegmatites: selected examples of data and mechanisms. *Canadian Mineralogist*, 23: 381–421.
- Claiborne, L.L.; Miller, C.F.; Walker, B.A.; Wooden, J.L.; Mazdab, F.K. & Bea, F. (2006). Tracking magmatic processes through Zr/Hf ratios in rocks and Hf and Ti zoning in zircons: An example from the Spirit Mountain batholith, Nevada. *Mineralogical Magazine*, 70: 517–543. <http://dx.doi.org/10.1180/0026461067050348>.
- Claiborne L.L.; Miller C.F. & Wooden J.L. (2010). Trace element composition of igneous zircon: A thermal and compositional record of the accumulation and evolution of a large silicic batholith, Spirit Mountain, Nevada. *Contributions to Mineralogy and Petrology*, 160: 511–31. <http://dx.doi.org/10.1007/s00410-010-0491-5>.
- Dias, G.; Simões, P.P.; Ferreira, N. & Leterrier, J. (2002). Mantle and crustal sources in the genesis of late-Hercynian granitoids (NW Portugal): geochemical and Sr-Nd isotopic constraints. *Gondwana Research*, 5: 287–305. [http://dx.doi.org/10.1016/S1342-937X\(05\)70724-3](http://dx.doi.org/10.1016/S1342-937X(05)70724-3).
- Ferreira, N.; Iglesias, M.; Noronha, F.; Pereira, E.; Ribeiro, A. & Ribeiro, M.L. (1987). Granitos da Zona Centro Ibérica e seu enquadramento geodinâmico. In: Geología de los Granitoides y Rocas Asociadas del Macizo Hespérico (Bea, F.; Carnicer, A.; Gonzalo, J.; López Plaza, M. & Rodríguez Alonso M., Eds), Editorial Rueda, Madrid, 37–51.

- Gagnevin, D.; Daly, J.S. & Kronz, A., (2010). Zircon texture and chemical composition as a guide to magmatic processes and mixing in a granitic environment and coeval volcanic system. Contributions to Mineralogy and Petrology, 159: 579–596. <http://dx.doi.org/10.1007/s00410-009-0443-0>.
- Hanchar, J.M. & Miller, C.F. (1993). Zircon zonation patterns as revealed by cathodoluminescence and backscattered electron images: implications for interpretation of complex crustal histories. Chemical Geology, 110: 1–13. [http://dx.doi.org/10.1016/0009-2541\(93\)90244-D](http://dx.doi.org/10.1016/0009-2541(93)90244-D).
- Heaman, L.M.; Bowins, R. & Crocket, J. (1990). The chemical composition of igneous zircon suites: implications for geochemical tracer studies. Geochimica et Cosmochimica Acta, 54: 1597–1607. [http://dx.doi.org/10.1016/0016-7037\(90\)90394-Z](http://dx.doi.org/10.1016/0016-7037(90)90394-Z).
- Hoskin, P.W.O.; Kiny, P.D.; Wyborn, D. & Chappell, B.W. (2000). Identifying accessory mineral saturation during differentiation in granitoid magmas: an integrated approach. Journal of Petrology, 41: 1365–1369. <http://dx.doi.org/10.1093/petrology/41.9.1365>.
- Hoskin, P. & Schaltegger U. (2003). The composition of zircon and igneous and metamorphic petrogenesis. Reviews in Mineralogy and Geochemistry, 53: 27–62. <http://dx.doi.org/10.2113/0530027>.
- Martins, H.C.B.; Sant’Ovaia, H. & Noronha, F. (2007). Post-tectonic granite intrusion controlled by a deep Variscan fault in Northern Portugal. Cadernos do Laboratorio Xeológico de Laxe, 32: 221–235.
- Martins, H.C.B.; Sant’Ovaia, H. & Noronha, F. (2009). Genesis and emplacement of felsic Variscan plutons within a deep crustal lineation, the Penacova-Régua-Verín fault: an integrated geophysics and geochemical study (NW Iberian Peninsula). Lithos, 111: 142–155. <http://dx.doi.org/10.1016/j.lithos.2008.10.018>.
- Martins, H.C.B.; Sant’Ovaia, H.; Abreu, J.; Oliveira, M. & Noronha, F. (2011). Emplacement of the Lavadores granite (NW Portugal): U/Pb and AMS results. Comptes Rendus Geoscience, 343: 387–396. <http://dx.doi.org/10.1016/j.crte.2011.05.002>.
- Martins, H.C.B.; Sant’Ovaia, H. & Noronha, F. (2013). Late-Variscan emplacement and genesis of the Vieira do Minho composite pluton, Central Iberian Zone: constraints from U-Pb zircon geochronology, AMS data and Sr-Nd-O isotope geochemistry. Lithos, 162–163: 221–235. <http://dx.doi.org/10.1016/j.lithos.2013.01.001>.
- Martins, H.C.B.; Simões, P.P. & Abreu, J. (2014). Zircon crystal morphology and internal structures as a tool for constraining magma sources: Examples from northern Portugal Variscan biotite-rich granite plutons. Comptes Rendus Geoscience, 346: 233–243. <http://dx.doi.org/10.1016/j.crte.2014.07.004>.
- Pérez-Soba, C.; Villaseca, C.; González del Tánago, J. & Nasdala, L. (2007). The composition of zircon in the peraluminous Hercynian granites of the Spanish Central System batholith. The Canadian Mineralogist, 4: 509–527. <http://dx.doi.org/10.2113/gscanmin.45.3.509>.
- Pupin, J.P. (2000). Granite genesis related to geodynamics from Hf-Y in zircon. Transaction of the Royal Society of Edinburg: Earth Sciences, 91: 254–256. <http://dx.doi.org/10.1017/S0263593300007410>.
- Silva, M.M.V.G. (1995). Mineralogia, petrologia, e geoquímica de encraves de rochas graníticas de algumas regiões Portuguesas. PhD thesis, Coimbra University, Portugal, 288 pp.
- Silva, M.M.V.G. & Neiva, A.M. (1998). Geoquímica de encraves microgranulares e granitos hospedeiros da região de Vila Nova de Gaia, Norte de Portugal. Comunicações do Instituto Geológico e Mineiro, 84: 35–38.
- Silva, M.M.V.G. (2010). O granito de Lavadores e os seus encraves. In: Ciências Geológicas – Ensino e Investigação e sua História (Cotelo Neiva, J.M.; Ribeiro, A.; Mendes Victor, L.; Noronha, F. & Magalhães Ramalho, M., Eds), 1: 269–279.
- Smith, D.G.W.; de St. Jorre, L.; Reed, S.J.B. & Long, J.V.P. (1991). Zonally metamictized and other zircons from Thor Lake, Northwest Territories. The Canadian Mineralogist, 29: 301–309.
- Speer, J.A. (1980). Zircon. In Orthosilicates (Ribbe, P.H. Ed). Reviews in Mineralogy, 5: 67–112.
- Watson, E.B. (1996). Dissolution, growth and survival of zircons during crustal fusion: kinetic principles, geological models and implications for isotopic inheritance. Transaction of the Royal Society of Edinburg: Earth Sciences, 87: 43–56. <http://dx.doi.org/10.1017/S0263593300006465>.
- Watson, E.B. & Cherniak D.J. (1997). Oxygen diffusion in zircon. Earth and Planetary Science Letters, 148: 527–544. [http://dx.doi.org/10.1016/S0012-821X\(97\)00057-5](http://dx.doi.org/10.1016/S0012-821X(97)00057-5).
- Watson, E.B., Wark, D.A. & Thomas, J.B. (2006). Crystallisation thermometers for zircon and rutile. Contributions to Mineralogy and Petrology, 151: 413–433. <http://dx.doi.org/10.1007/s00410-006-0068-5>.